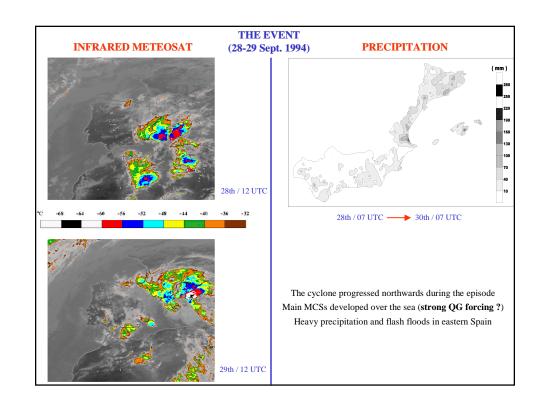
## ESTUDIO NUMÉRICO DE LA PREDICTABILIDAD DE UN EVENTO DE CICLOGÉNESIS MEDITERRÁNEA MEDIANTE INVERSIÓN DE VORTICIDAD POTENCIAL

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#### CONTROL NUMERICAL SIMULATION

\* PSU-NCAR mesoscale model (non-hydrostatic version MM5)

#### \* Simulation:

- 2 domains: 82x82x31 (60 and 20 km)

- Interaction: two-way

- I.C and B.C: NCEP global analysis + Surface and Upper air obs.

- Period: 48 h, from 00 UTC 28 September 1994

#### \* Physical parameterizations:

- PBL: Based on Blackadar (1979) scheme (Zhang and Anthes 1982)

- Ground temperature: Force-restore slab model (Blackadar 1979)

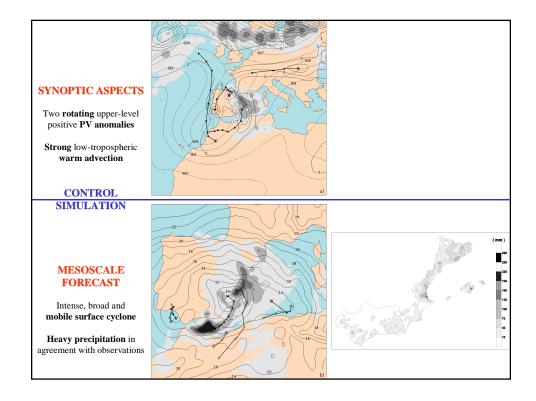
- Radiation fluxes: Considering cloud cover (Benjamin 1983)

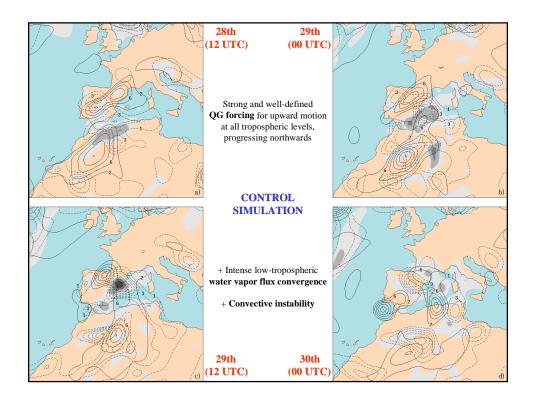
- Resolved-scale microphysics:

Cloud water, rainwater, cloud ice and snow (Dudhia 1989)

- Parameterized convection:

60 km: Betts-Miller (1986) 20 km: Kain-Fritsh (1990)





# SENSITIVITY TO THE UPPER LEVEL PV ANOMALIES (motivation)

- \* The two embedded upper-level PV centres seem to be playing an important role for the evolution, intensity and areal extent of the surface cyclone
- \* How a potential analisis and/or forecast error in the representation of these PV anomalies might affect the mesoscale forecast ?



- \* Sensitivity analysis based on additional simulations with perturbed initial conditions
- \* A balanced flow associated with each anomaly must be found that can be used to alter the model initial conditions in a physically consistent way without introducing any significant noise in the model ———> Piecewise PV inversion

## PIECEWISE PV INVERSION TECHNIQUE (Davis and Emanuel; MWR 1991)

- 1) Balanced flow  $(\phi, \psi)$  given instantaneous distribution of Ertel's PV (q):
- \* Charney (1955) nonlinear balance equation

$$\nabla^2 \phi = \boldsymbol{\nabla} \cdot f \boldsymbol{\nabla} \psi + 2m^2 \left[ \frac{\partial^2 \psi}{\partial x^2} \frac{\partial^2 \psi}{\partial y^2} - \left( \frac{\partial^2 \psi}{\partial x \partial y} \right)^2 \right]$$

f Coriolis parameter

m map-scale factor

\* Approximate form of Ertel's PV 
$$q = \frac{g\kappa\pi}{p} \left[ (f + m^2 \nabla^2 \psi) \frac{\partial^2 \phi}{\partial \pi^2} - m^2 \left( \frac{\partial^2 \psi}{\partial x \partial \pi} \frac{\partial^2 \phi}{\partial x \partial \pi} + \frac{\partial^2 \psi}{\partial y \partial \pi} \frac{\partial^2 \phi}{\partial y \partial \pi} \right) \right]$$

$$p \text{ pressure} \qquad g \text{ gravity} \qquad \kappa = Rd/Cp \qquad \pi = Cp(p/po)^{\kappa}$$

- \* **Bounday conditions** Lateral (Dirichlet) / Top and Bottom(Neumann):  $\partial \phi/\partial \pi = f \partial \psi/\partial \pi = -\theta$  $\theta$  potential temperature
- 2) Reference state: Balanced flow  $(\bar{\phi}, \bar{\psi})$  given time mean distribution of Ertel's PV  $(\bar{q})$ :
- \* Same equations as in 1), except using time mean fields instead of instantaneous fields
- 3) Perturbation fields ( $\phi$ ',  $\psi$ ', q') given by the definitions:  $(q, \phi, \psi) = (\bar{q}, \bar{\phi}, \bar{\psi}) + (q', \phi', \psi')$

#### PIECEWISE PV INVERSION TECHNIQUE

- 4) We consider that q' is partitioned into N portions or anomalies:  $q' = \sum_{n=1}^{N} q_n$
- 5) <u>Piecewise inversion</u>:  $(\phi_n, \psi_n)$  associated with  $q_n$ ? ... and requiring:  $\psi' = \sum_{n=1}^N \phi_n$  $\psi' = \sum_{n=1}^N \phi_n$

. and requiring: 
$$\phi' = \sum_{n=1}^N \phi_n$$

...After substitution of the above summations in the balance and PV equations and some rearrangements

$$\nabla^2 \phi_n = \boldsymbol{\nabla} \cdot f \boldsymbol{\nabla} \psi_n + 2m^2 \left( \frac{\partial^2 \psi^*}{\partial x^2} \frac{\partial^2 \psi_n}{\partial y^2} + \frac{\partial^2 \psi^*}{\partial y^2} \frac{\partial^2 \psi_n}{\partial x^2} - 2 \frac{\partial^2 \psi^*}{\partial x \partial y} \frac{\partial^2 \psi_n}{\partial y \partial x} \right)$$

$$\begin{array}{ll} q_n & = & \displaystyle \frac{g\kappa\pi}{p} \left[ (f + m^2\nabla^2\psi^*) \frac{\partial^2\phi_n}{\partial\pi^2} + m^2 \frac{\partial^2\phi^*}{\partial\pi^2} \nabla^2\psi_n \right. \\ & \left. \left. - m^2 \left( \frac{\partial^2\phi^*}{\partial x\partial\pi} \frac{\partial^2\psi_n}{\partial x\partial\pi} + \frac{\partial^2\phi^*}{\partial y\partial\pi} \frac{\partial^2\psi_n}{\partial y\partial\pi} \right) - m^2 \left( \frac{\partial^2\psi^*}{\partial x\partial\pi} \frac{\partial^2\phi_n}{\partial x\partial\pi} + \frac{\partial^2\psi^*}{\partial y\partial\pi} \frac{\partial^2\phi_n}{\partial y\partial\pi} \right) \right] \end{array}$$

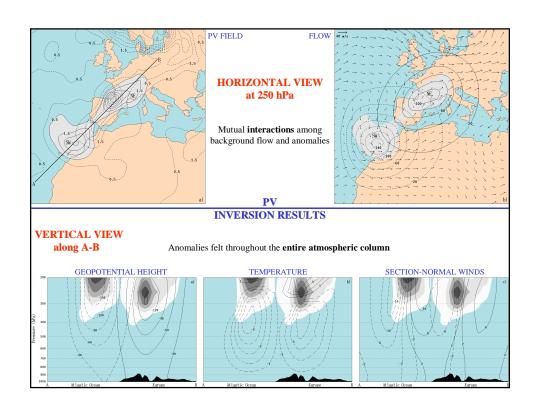
where  $()^* = \overline{()} + \frac{1}{2}()'$ 

**Boundary conditions:** Lateral (homogeneous) / Top and bottom (using  $\theta_n$ )

At 00 UTC 28 September 1994, using the NCEP-based isobaric analysis

\* In our case study: Reference state: 6-day time average about 00 UTC 28 September

Anomalies: positive PV perturbations above 500 hPa SW and NE of Gulf of Cádiz



## SENSITIVITY EXPERIMENTS

By adding and/or subtracting the PV-inverted balanced fields (geopotential, temperature and wind) into the model initial conditions

#### Sensitivity to the intensity

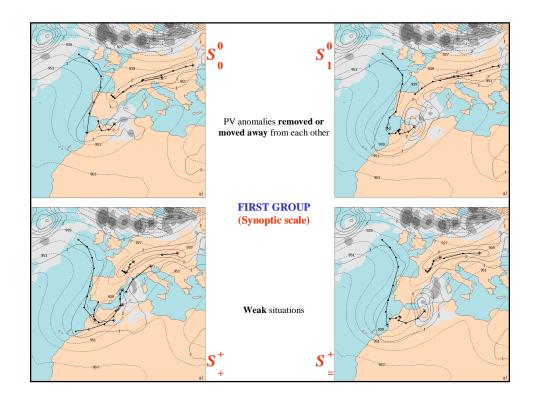
(One or both PV anomalies removed or doubled)

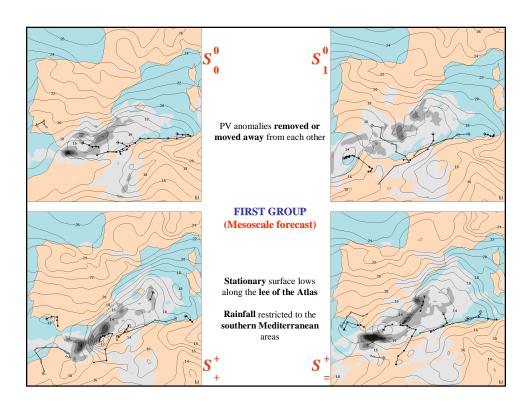
### Sensitivity to the position

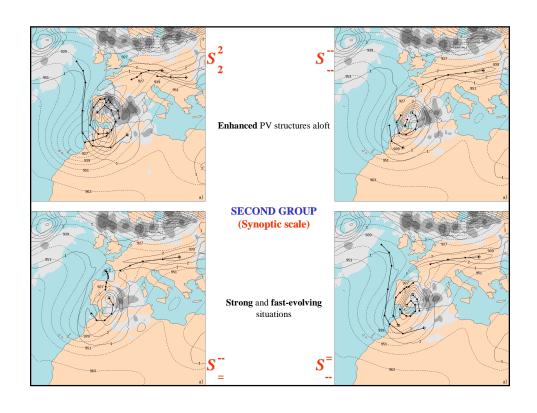
(One or both PV anomalies shifted 425 km along A-B)

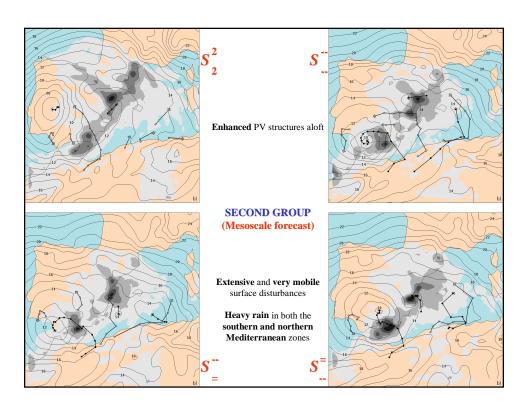
Experiment	SW anomaly	NE anomaly
$S_0^0$	Removed	Removed
$S_2^2$	Doubled	Doubled
$S_1^0$	Unchanged	Removed
$S_2^0$	Doubled	Removed
$S_0^1$	Removed	Unchanged
$S_0^2$	Removed	Doubled
$S_2^1$	Doubled	Unchanged
$S_{1}^{2}$	Unchanged	Doubled

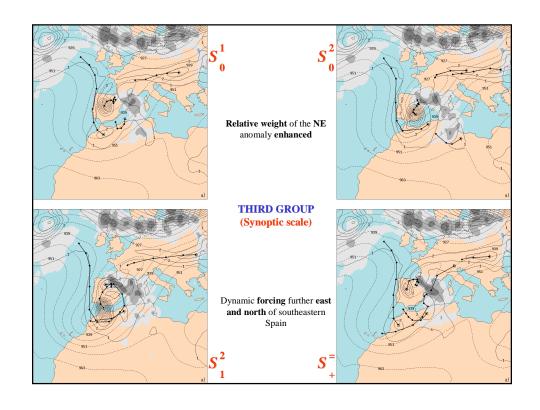
Experiment	SW anomaly	NE anomaly
S_	Moved inwards	Moved inwards
S <sup>+</sup> _+	Moved outwards	Moved outwards
S=	${\bf Unchanged}$	Moved inwards
S <sub>+</sub>	Moved outwards	Moved inwards
S=	Moved inwards	Unchanged
S_+	Moved inwards	Moved outwards
S=	Moved outwards	Unchanged
S <sub>=</sub> <sup>+</sup>	${\bf Unchanged}$	Moved outwards

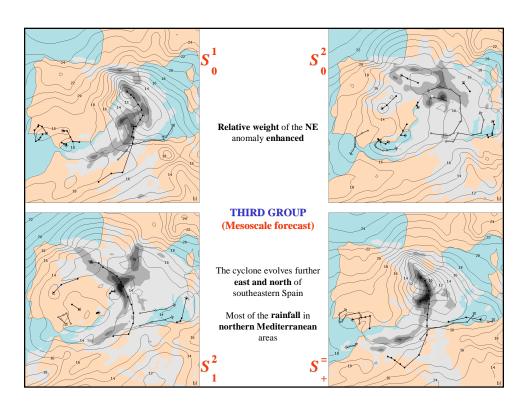


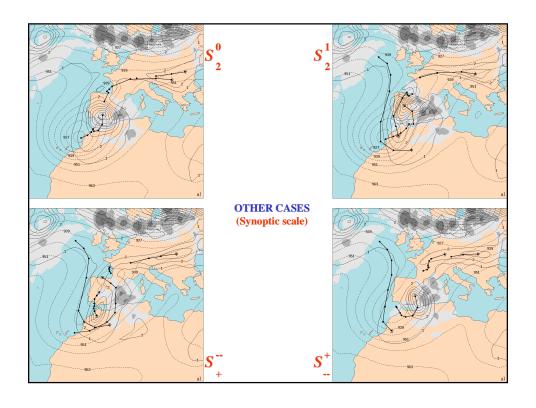


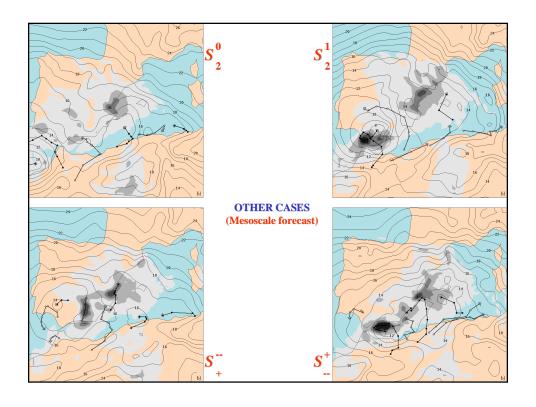


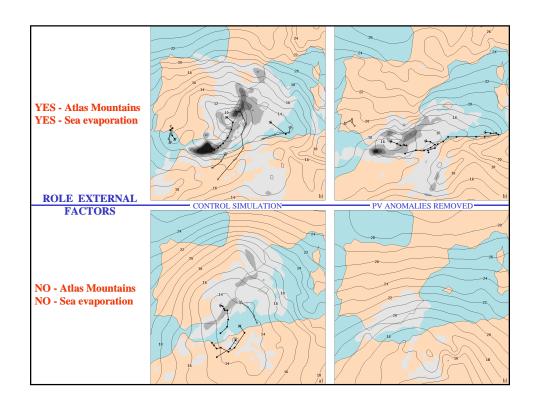












## **CONCLUSIONS**

- \* Track, shape and intensity of the surface cyclone and the corresponding rainfall pattern are very sensitive to the embedded upper-level PV anomalies (a potential error in the initial representation of the anomalies can be critical)
- \* The external factors induced an appreciable modulation of the surface circulation and enhanced the efficiency of the system as a rainfall producer, but the cyclogenesis over the southern Mediterranean and its progression to the north must be attributed mostly to the action of the upper-level PV anomalies
- \* The combined application of piecewise PV inversion and numerical simulation offers a valuable and unique framework from which the effects of dynamical features of the flow can be studied in a practical and physically consistent way